

INFLUENCE OF THE LARGE CATION UPON PRE-EDGE DETAILS OF Fe 1s XANES SPECTRA IN JAROSITE-TYPE SULPHATE MINERALS

M.O. FIGUEIREDO & T. PEREIRA da SILVA

Cryst, Miner, Centre, IICT & INETI/IGM, Geological Data Centre, Estrada da Portela, 2721-866 Alfragide and CENIMAT / I3 N, Mater. Sci. Dpt., Fac. Sci. & Technology, New Univ. Lisbon, 2829-516 Caparica, Portugal



The jarosite-alunite mineral group has the simplified *crystal*

chemical formula Aic Bo3(Tt O4)2(OH)6 where ic stands

for *pseudo-icosahedral* (coordination number CN = 12, with $\underline{A} = K^+$, Na⁺, NH₄⁺, Ag⁺, Tl⁺, Pb²⁺, Bi³⁺ or even H₃O⁺, and minor

Ca2+, Ba2+, Sr2+ plus trivalent rare-earth ions, namely Eu3+), o

for octahedral (CN=6, with B standing for Fe3+ (jarosite s.s.) or

for Al3+ (alunite s.s.) and also for V3+, plus minor Fe2+ along

with other divalent cations, Cu2+, Zn2+, Mg2+), t for tetrahedral

Octahedral B cations are coordinated by 4OH and 2O-anions

from tetrahedral groups, while A cations are surrounded by

6OH shared with B-cation octahedra plus 6O-atoms from the

Jarosite crystal structure (fig.1) was determined in 1937 [3]. It is

trigonal, space group $R \overline{3} m$, and the unit cell (hexagonal

description, with Z=3 and c/a ratio ~2.5) contains 3A and

9 B atoms occupying invariant equipoints and 6 T plus

6O-atoms filling mono-variant sites (positional free parameter

z along the 3-fold axis) and 18O-atoms plus 18 OH replaced by

minor H2O in bi-variant sites with x,z free parameters. Na-

jarosite and Pb-jarosite are isostructural and form solid solution

with jarosite despite the different nature of the A-cation.

tetrahedral groups (including the non-shared oxygens).

Crystallography of Jarosites

(CN = 4, with $\underline{T} = S^{6+}, P^{5+}, As^{5+}$).

Problematic

Jarosite s.s. - ideally KFe3(OH)6(SO4)2 - is a secondary iron sulphate very common in acid mine drainage (AMD) environments with high potassium content.

This mineral is not only environmentally important because it may configure the ultimate mineral sink of lead and other hazardous metals (e.g. thallium), beyond silver, occurring frequently in the oxidized part of sulfide ore deposits, but it is also economically relevant within various hydrometallurgical routes.

Besides, the 2D Kagomé-pattern of iron octahedra in the crystal structure (fig.1) awards unique magnetic properties to this mineral. The octahedral crystal field stabilizes high spin-state ferric ions, and this peculiar undistorted triangular motif - with fully occupied magnetic sites and weak interlayer coupling - renders jarosite an important, paradigmatic Kagomé Heisenberg Anti-Ferromagnet, KHAFM [1]. Remarkably, jarosite is also the terrestrial analogue of Martian sulphates [2].

Once many doubts still subsist concerning the exact crystal structure of Pb-rich and other jarosites, a comparative study on the electronic state of iron in these important sulphates was undertaken by comparing and analysing the pre-edge details of Fe 1s XANES spectra from various jarosites and structurally afine iron sulphates (Table 1).



Experimental

X-ray absorption experiments were carried out at the ESRF using beam line ID-21 [6]. XANES spectra were collected in fluorescence yield (FY) mode with an energy-dispersive high-purity Gedetector mounted perpendicular to the X-ray beam in the horizontal plane (fig.4) by irradiating a small area (~1µm2) of well crystallized mineral fragments.

A fixed-exit Si (220) monochromator assuring an energy resolution of 0,3 eV was used for the energy scans (7050-7350 eV). Metallic iron was used for energy calibration.

Pre-edge details of XANES spectra were analysed using Fityk program [7] to deconvolute this spectral region into pseudo-Voigt components and assess iron electronic state.

* The support from EU through the Action Access to Research Infrastructures is acknowledged. Special thanks are due to the ESRF beamline Scientists, Drs. J. Susini & V. de Andrade, and to our Colleagues J. Veiga & J. Mirão for participating in data collection through the experiment EC-87.



[S'04] [Fe°X₆] Studied sulphate minerals The number of iron coordination polyhedra (octahedra, o) (H2O)3] [Fe ° O4 (OH) 2] (H2O)6] 0202 ^t0<u>0</u>3] 0°] tion polyhedra (tetrahedra, t) in the and sulphur coordin solated unit cell of the mineral is indicated [Fe°O3 Fe° ô Oxygen anions shared between Fe & S are underlined [Fe ŝ S 04] ²H² Chemical Formula [crist struct.] Name ŝ K Fe 3 (OH)6 (SO4)2 [fig. 1] JAROSITE 3 2 Fe^{II} (SO₄).7 H₂O [fig. 2] MELANTERITE 1 1 1 COQUIMBITE Fe^{III}₂ (SO₄)₃.9 H₂O [fig. 3] 3 1 1 1 6

Table 1

Plumbojarosite: accommodation of the lone pair of 5s² electrons from Pb2+ ions without symmetry breaking



Results and conclusions

The pre-edge features of Fe 1s XANES spectra collected from studied sulphate minerals are reproduced in fig.5. As expected [8], the average pre-peak centroid (fig. 5a) clearly differentiates between ferrous and ferric states: 7112.5 vs. 7114 eV, close to the published value of 1.4 eV for the energy difference [9]. The similitude of the main pre-edge contributions is remarkable despite the occurrence of three distinct Fe3+ coordination environments in coquimbite comparatively to one unique situation in jarosite (Table 1).

Conversely, the size and speciation or electronic state of the large cation -in icosahedral coordination by 12 O-atoms from the sulphate tetrahedral groups may explain the observed differences (fig.5b) : (1) K+ and Na+ are closed shell, non-polarisable alkaline ions with different radius inducing a smaller c/a ratio for natrojarosite; (2) off-centering of Pb2+ ions within the icosahedron (fig. 6) to accommodate the lone-pair of 5s² electrons through a statistical partial occupation of a non-invariant equipoint within the same space group symmetry.





(b)

K +

Ph 2+

Na +



References

- [1] R. BALLOU, et al. (2003) J. Magnetism [7] FITYK program: & Magnet. Mat. 262, 465.
- [2] J.J. PAPIKE, J.M. KARNER & C.K. SHEARER (2006) Geochim. Cosmoch. Acta 70, 1309.
- [3] B. HENDRICKS (1937) Amer. Inst. Min. Met. Publ. 22, 773.
- [4] W.H. BAUR (1964) Acta Cryst. 17, 1167 [5] J.H. FANG & P.D. ROBINSON (1970)
- Amer. Min. 55, 1534
- [6] J. SUSINI, et al. (2002) Surface Rev and Letters 9, 203.

- http://www.unipress.waw.pl/fityk/ [8] T.E. WESTRE. et al. (1997) J.
- Amer. Chem. Soc. 119, 6297. [9] M. WILKE, et al. (2001) Amer.
- Min. 86, 714. [10] J.T. SZYMANSKI (1985) Canad.
- Min. 23, 659. [11] I.T. SZYMANSKI (1988) Canad
- Min. 26, 923. [12] L.C. BASCIANO & R.C.
 - PETERSON (2008) Amer.Min. 93. 853